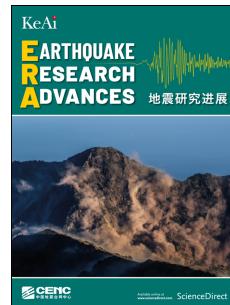


# Journal Pre-proof



Characteristics of Ambient Seismic Noise Recorded at Offshore Wind Turbine Platform Monitoring Stations

Jun Wang, Jiang-rong Zheng, Ye-jun Sun, Ling-yu Xu, Yi-cheng He, Jie Gong, Chen Wang, Xiao-yan Zhan, Yu-jie Wan, Wen-jie Ren

PII: S2772-4670(25)00083-1

DOI: <https://doi.org/10.1016/j.eqrea.2025.100440>

Reference: EQREA 100440

To appear in: *Earthquake Research Advances*

Received Date: 3 April 2025

Revised Date: 8 November 2025

Accepted Date: 24 November 2025

Please cite this article as: Wang, J., Zheng, J.-r., Sun, Y.-j., Xu, L.-y., He, Y.-c., Gong, J., Wang, C., Zhan, X.-y., Wan, Y.-j., Ren, W.-j., Characteristics of Ambient Seismic Noise Recorded at Offshore Wind Turbine Platform Monitoring Stations, *Earthquake Research Advances*, <https://doi.org/10.1016/j.eqrea.2025.100440>.

This is a PDF of an article that has undergone enhancements after acceptance, such as the addition of a cover page and metadata, and formatting for readability. This version will undergo additional copyediting, typesetting and review before it is published in its final form. As such, this version is no longer the Accepted Manuscript, but it is not yet the definitive Version of Record; we are providing this early version to give early visibility of the article. Please note that Elsevier's sharing policy for the Published Journal Article applies to this version, see: <https://www.elsevier.com/about/policies-and-standards/sharing#4-published-journal-article>. Please also note that, during the production process, errors may be discovered which could affect the content, and all legal disclaimers that apply to the journal pertain.

© 2025 China Medical Cosmetology Press Co. Ltd. Publishing services by Elsevier B.V. on behalf of KeAi Communications Co. Ltd.

# Characteristics of Ambient Seismic Noise Recorded at Offshore Wind Turbine Platform Monitoring Stations

Jun Wang<sup>1</sup>, Jiang-rong Zheng<sup>1</sup>, Ye-jun Sun<sup>1</sup>, Ling-yu Xu<sup>2</sup>, Yi-cheng He<sup>1</sup>,

Chen Wang<sup>1</sup>, Jie Gong<sup>1</sup>, Yu-jie Wan<sup>3</sup>, Wen-jie Ren<sup>2</sup>

<sup>1</sup>Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.  
Email: wangjun1099@qq.com, Phone: 13951702501.

<sup>2</sup>Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.  
Email: jsdzjzjr@126.com.

<sup>3</sup>Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.  
Email: syj\_4116@126.com.

<sup>4</sup>Professor, Institute of Geotechnical Engineering, Nanjing Tech University, Nanjing 210009, China. Email: lyxu@njtech.edu.cn

<sup>5</sup>Associate professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China. Email: hyckevin@mail.ustc.edu.cn

<sup>6</sup>Engineer, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.  
Email: 372766770@qq.com

<sup>7</sup>Associate professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China. Email: 498602259@qq.com

<sup>8</sup>Engineer, Earthquake Administration of Qinghai Province, Xining 810000, China.  
Email: 376633212@qq.com

<sup>9</sup>PhD candidate, Institute of Geotechnical Engineering, Nanjing Tech University, Nanjing 210009, China. Email: renwenjie@njtech.edu.cn

**Corresponding author:** Jun Wang, wangjun1099@qq.com

## Characteristics of Ambient Seismic Noise Recorded at Offshore Wind Turbine Platform Monitoring Stations

**Abstract:** This study examines ambient seismic noise recorded from operational seismic monitoring stations installed on offshore wind platforms in the Yellow Sea. The research utilizes one-year three-component continuous waveform data to investigate energy intensity, wavefield composition, and polarization properties through frequency-domain polarization analysis. A dynamic finite element analysis is conducted on a typical offshore platform structure to investigate the amplification effect. The results show that: (1) The energy of single-frequency microseisms is clearly observable, while distinct segmentation phenomena are observed near 0.2 Hz within the double-frequency microseism (DF) band, with short-period DF exhibiting stronger energy than long-period DF. The wind, wave and current may result in greater horizontal noise energy intensity than vertical components at specific frequencies and directions; (2) The ambient seismic noise recorded at offshore platform monitoring stations exhibits systematic amplification compared to onshore station observations, with an average amplification factor of 3-5 across the studied frequency band. Notably, maximum amplification reaches 6.6-7.7 times within the 1.2-1.6 Hz range, representing a significant resonant response characteristic of the offshore platform structures. (3) Within the microseismic band (20 s-0.5 Hz), the azimuth of the noise polarization principal axis predominantly clusters around 200°. The polarization degree exhibits perturbed variations with frequency between 0.2-0.6 Hz and is slightly greater than that obtained from onshore stations. These observations indicate that hurricane/storm activities and short-period ocean waves in the relevant maritime area, along with their breaking and turbulent processes, generate pronounced high-frequency noise components.

26 **Keywords:** Marine seismic observation; Offshore wind power platform; Ambient seismic noise;  
27 Frequency-domain polarization analysis; Finite element modeling

28 1. Introduction

29 Seismic monitoring provides a basis for understanding earthquake preparation and occurrence  
30 mechanisms, investigating the Earth's internal structure, and implementing earthquake early  
31 warning systems. However, the current seismic monitoring relies on land-based seismic networks  
32 due to the challenges posed by the ocean environment, resulting in a sparse distribution of marine  
33 seismic observation stations. This limitation is particularly alarming given that oceans cover  
34 approximately 71% of the Earth's surface and marine earthquakes account for about 85% of global  
35 seismic activity, as evidenced by global seismic data. With the rapid development of the marine

36 economy, the need for marine seismic observations and the mitigation of marine seismic hazards  
37 has become increasingly urgent (Chen et al., 2025; Xu et al., 2025b). Consequently, since the 1960s,  
38 many countries have initiated marine seismic observation efforts, progressively advancing research  
39 in this field and contributing significantly to marine seismic monitoring, marine seismic structure  
40 surveys, and investigation of secondary marine disasters (Stephen, 2003).

41 With the advancement of science and technology, seismologists have been continuously  
42 striving to apply various advanced seismic observation technologies, such as ocean-bottom  
43 seismometers (OBS), cabled systems, buoy-based systems, and distributed fiber-optic sensing  
44 (DAS), to marine seismic monitoring. However, significant challenges in power supply, data  
45 communication, and long-term maintenance, preventing the implementation of sustained,  
46 continuous, and fixed seismic observations. These limitations have severely hindered substantial  
47 progress in marine seismic observation and its applied research. Thus, most research remain focused  
48 on technical validations and short-term seismic recordings. For example, Ito et al. (2017) conducted  
49 a one-year observation of volcanic activity in the Pacific using OBS to investigate the earthquake  
50 localization and velocity structure. Lin et al. (2024) captured global seismic data during a 3- to 6-  
51 month deployment of broadband OBS, including the 7.8 magnitude earthquake in New Zealand  
52 and the 6.3 magnitude nuclear test in North Korea. Krylov et al. (2021) reported the effectiveness  
53 of broadband Molecular Electronic Transfer (MET) sensors in OBS for high-quality seismic data  
54 under Arctic conditions, highlighting their potential use for seismic hazard assessments in Arctic  
55 region. In addition, the Earthquake Research Institute of the University of Tokyo has deployed a  
56 submarine earthquake and tsunami observation cable system using optical fiber for data  
57 transmission in the Sanriku region since 1996. This system has recorded small seismic events of  
58 around magnitude 1.8 near the submarine cable and a distant earthquake of magnitude 6.6 with an  
59 epicenter approximately 2,300 kilometers away (Shinohara et al., 2022). Baba et al. (2024)  
60 conducted marine seismic observation by deploying Distributed Acoustic Sensing (DAS) in the  
61 Tsugaru Strait and detected earthquakes with magnitudes below 1.0 during a four-month  
62 observation period.

63 In recent years, the rapid growth of the offshore wind power (OWP) industry has provided  
64 new opportunities for marine seismic observation due to the increasing global demand for  
65 renewable energy. OWP platforms offer a promising foundation for constructing seismic  
66 monitoring stations because they are widely distributed and offer unique advantages in power

67 supply and communication. Since December 2022, the Earthquake Administration of Jiangsu  
68 Province (EAJ) of China has progressively established seismic monitoring stations on the OWP  
69 platforms (Sun et al., 2022; Zhu et al., 2023) and successfully recorded the dynamic response of  
70 the OWP platforms under earthquakes in the Yellow Sea region. In September 2023, the China  
71 Earthquake Administration released a document titled "China Marine Seismic Observation Plan  
72 (2023-2035)", explicitly proposing the construction of seismic monitoring stations on OWP  
73 platforms. Consequently, it is anticipated that more seismic monitoring stations will be deployed in  
74 the future. As an effective approach to marine seismic observation, the extraction of spectral  
75 characteristics from recorded data and their subsequent application to seismic research and  
76 earthquake early warning systems in maritime regions constitute critical scientific questions in the  
77 field of seismic monitoring. A comprehensive analysis of ambient seismic noise in continuous  
78 recordings from offshore platform monitoring stations offers valuable insights for assessing data  
79 quality and utility.

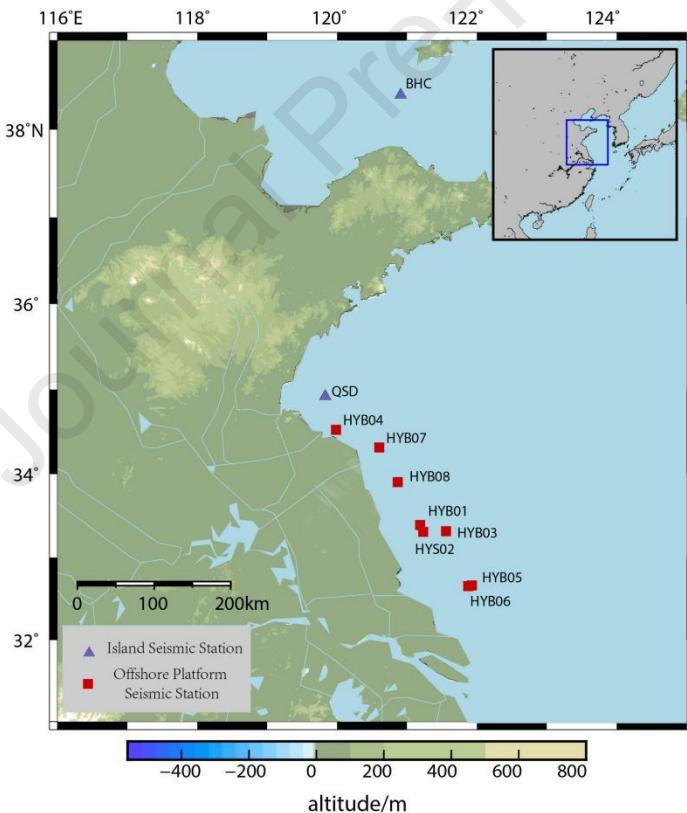
80 In seismological research, microseism is defined as the most energetic component of ambient  
81 seismic noise within the 2-20 second period band. This band exhibits two distinct spectral peaks,  
82 classified as single-frequency (SF) microseisms and double-frequency (DF) microseisms, each  
83 generated by different excitation mechanisms (Miche, 1944; Hasselmann, 1963; Bromirski et al.,  
84 2005; Tanimoto & Prindle, 2007; Koper & Burlacu, 2015; Xiao et al., 2018; Wang Jun et al., 2022).  
85 Thus, this study examines continuous waveform data recorded by EAJ between May 2023 and  
86 December 2024. The analytical framework incorporates two advanced methodologies: (1)  
87 frequency-domain polarization analysis and (2) dynamic response analysis of OWP platforms. This  
88 study systematically investigates three fundamental aspects of microseismic phenomena: (1) the  
89 source characteristics and underlying generation mechanisms of microseismic noise, (2) the spectral  
90 properties of ambient seismic noise, and (3) the potential influence of offshore platform structures  
91 on microseismic noise. The results provide critical foundations for developing noise seismology  
92 applications and optimizing the use of offshore monitoring station data in seismic research.

## 93 2. Offshore platform monitoring stations

94 Over the years, marine seismic observation has faced persistent operational limitations due to  
95 power supply constraints, unreliable communications, and maintenance difficulties, preventing  
96 sustained, high-quality seismic monitoring. The integration of seismic monitoring stations with

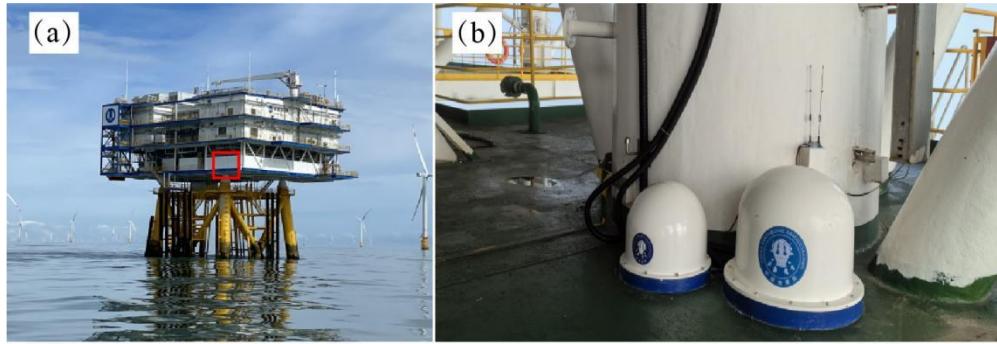
97 offshore wind platforms now offers a viable solution to these challenges. These installations deploy  
 98 broadband seismometers that serve three critical functions: (1) enhancing earthquake detection and  
 99 early warning precision, (2) facilitating high-resolution imaging of marine crustal structures, and  
 100 (3) enabling structural health monitoring of offshore installations. The EAJ has established seven  
 101 seismic monitoring stations in the Yellow Sea area, utilizing offshore wind power substations and  
 102 high-resistance stations as deployment platforms. The distribution of these offshore platforms is  
 103 listed in Fig. 1. Each monitoring substation is equipped with broadband seismometers and  
 104 accelerometers, strategically positioned near the primary support columns on the first floor of the  
 105 platforms, as shown in Fig. 2. The instruments are sealed and protected using an RH-S700  
 106 protective housing. The observation data are transmitted in real-time to the data center of the EAJ  
 107 via a submarine communication system (Gong et al., 2023). Table 1 provides detailed information  
 108 regarding the platform dimensions, offshore distances, instrument models, and seawater depths for  
 109 each monitoring substation.

110



111

112 Fig. 1 Distribution of seven seismic monitoring stations on offshore wind power high-voltage substations  
 113 and step-up substations



114

115 Fig. 2 Offshore wind power platform and seismic monitoring devices: (a) Offshore wind  
 116 substation (red rectangle indicates the installation position for monitoring devices) and (b) A  
 117 photo showing the installation position of monitoring devices

118

Table 1 Detailed information of seven monitoring substation.

No.	Name of Substations	Station Code	Equipment	Distance from Shore (km)	Platform Dimensions
1	Dafeng Phase I Step-Up Substation H11	HYB01	GL-PCS60	45	36.50m × 42.00m, 4 floors, height 22m
2	Dafeng Phase II High-resistance Substation H8-2	HYS02	GL-PS2	51	47m × 29.8m, 4 floors, weight approx. 2800 tons
3	Xiangshui Step-Up Substation	HYB04	GL-PCS60	10	25m × 28m, height approx. 25m, 4 floors, total weight approx. 2000 tons
4	Rudong Step-Up Substation H6	HYB05	GL-PCS60	60	39.8m × 43.5m, height 16m
5	Rudong Step-Up Substation H10	HYB06	GL-PCS60	60	39.8m × 43.5m, height 16m, 3 floors
6	Binhai Step-Up Substation H3	HYB07	GL-PCS60	36	38.1m × 41.54m, height 20.9m, 5 floors, weight approx. 3030 tons
7	Sheyang Step-Up Substation H2	HYB08	GL-PCS60	45	56m × 50m, height 29m, weight 3850 tons

119

Table 2 Detailed parameters of two equipment used for marine seismic observation

Serial Number	Equipment	Seismometer Sensitivity	Data Logger Conversion Factor	Sampling Rate	Frequency Bandwidth
1	GL-PCS60	1000 V/m/s	74.5nV/count	100Hz	60s-50Hz
2	GL-PS2	1000 V/m/s	74.5nV/count	100Hz	2s-50Hz

120 3. Analysis Procedure

121 **3.1 Polarization Analysis**

122 Polarization analysis is a classic and highly effective technique for extracting kinematic  
 123 features and characteristics from different components of particle motion (Jurkevics, 1988). It can  
 124 be classified into two main approaches: time-domain algorithms (Flinn, 1965; Lin et al., 2012) and

frequency-domain algorithms (Park et al., 1987; Chen et al., 2007). This study employs a frequency-domain polarization algorithm, which offers the advantage of using prolate spheroidal wave functions for tapering when directly obtaining the signal spectrum, thereby preserving the integrity of the cross-spectral matrix with minimal distortion. The theoretical foundation of this method has been well-documented (Samson, 1983; Park et al., 1987) and will not be elaborated here. The processing of single-station data follows established methodologies (Wang et al., 2020; Wang et al., 2022). Our proposed method assumes that Rayleigh waves are incident as plane waves on each seismic station (Fig. 3). The three-component continuous waveform data are processed using 1-hour segments as computational samples, with a 50% overlap ratio applied to the sliding windows of the  $3 \times 3$  complex spectral covariance matrices. The spectral estimation employs 10 tapered sub-windows, each with a duration of 819.2 seconds, yielding effective frequency results in the range of 0.016-10 Hz. Prior to computation, the raw data undergoes necessary preprocessing including resampling to 20 Hz and removal of instrument responses.

The polarization analysis extracts the four key parameters:

(1) Power of the polarization corresponding to the largest eigenvalue ( $\lambda_0$ ): This represents the power of the principal polarization derived from the analysis of the original three components, expressed in units of acceleration power spectral density (dB,  $10 \times \log_{10}(m^2/s^4/Hz)$ ) (Jepsen and Kennett, 1990; Wagner and Owens, 1996). These three components are the vertical (UD), north-south (NS), and east-west (EW) directions. It is a function of azimuth and effectively characterizes the vibrational energy of the signal. Regardless of how the signal is polarized or divided among the three components, its amplitude can be accurately captured across the entire frequency range.

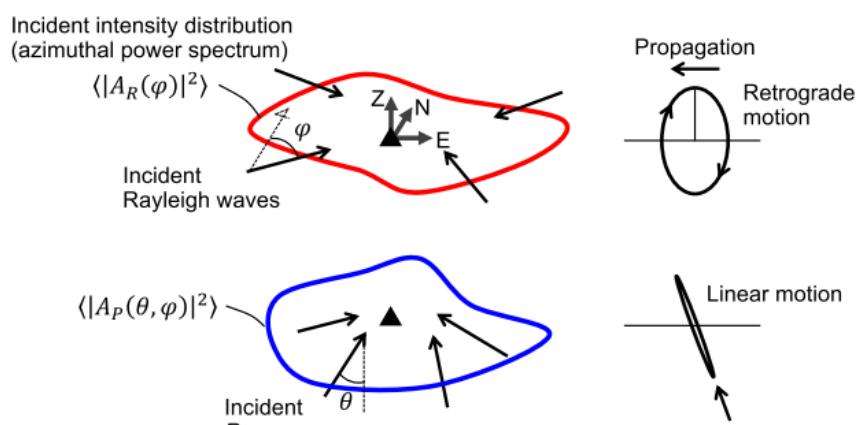


Fig. 3 Schematic illustration of Rayleigh waves and P waves incident on a single three-component station, and particle motions of fundamental-mode Rayleigh (R) and P waves. Black arrows indicate incident Rayleigh and P

149 waves.  $\varphi$  and  $\theta$  are propagation azimuth and incident angle. (Takagi et al., 2018)

150 (2) Degree of polarization ( $\beta^2$ ): This parameter is given in Eq. (1):

$$151 \quad \beta^2 = \frac{nTrS^2 - (TrS)^2}{(n-1)(TrS)^2} \quad (1)$$

152 where  $S$  is the spectral density matrix,  $n$  is the length of the data sequence,  $Tr$  is the minimization  
153 factor under non-negative constraints.  $\beta^2$  characterizes the degree of signal structure and is a  
154 dimensionless quantity representing the correlation between the components of the vibration vector.  
155  $\beta^2$  ranges between 0 and 1:  $\beta^2 = 0$  when all eigenvalues are equal and  $\beta^2 = 1$  when only one non-  
156 zero eigenvalue exists. For further details, refer to Eq. (31) in Samson (1983).

157 (3) Polarization ellipse azimuth ( $\theta_H$ ): This parameter represents the azimuth of the major axis  
158 of the polarization ellipse, ranging from  $0^\circ$  to  $360^\circ$ . It reflects the back-azimuth of wave sources  
159 such as Rayleigh waves or P-waves. Additionally,  $\theta_V$  represents the incidence angle of P-waves.

160 (4) Phase difference between radial and vertical components ( $\varphi_{VR}$ ): This parameter ranges  
161 from  $-90^\circ$  to  $90^\circ$ . There exists an equivalence effect between  $\theta_H$  and  $\varphi_{VR}$  (Koper and Burlacu,  
162 2015). For example, a retrograde Rayleigh wave ( $\varphi_{VR} = -90^\circ$ ) arriving from a specific back-  
163 azimuth ( $\theta_H = 45^\circ$ ) can be equivalently described as a prograde Rayleigh wave ( $\varphi_{VR} = 90^\circ$ ) arriving  
164 from the opposite direction ( $\theta_H = 225^\circ$ ).

## 165 3.2 Finite element modeling

166 To analyze the correlation between microseisms and platform structure response, this study  
167 uses the implicit dynamics module of the ABAQUS finite element software to develop a mechanical  
168 model of the offshore jacket platform. The governing equation for the implicit dynamics analysis  
169 module is:

$$170 \quad \mathbf{M}\ddot{\mathbf{u}} + \mathbf{C}\dot{\mathbf{u}} + \mathbf{K}\mathbf{u} = \mathbf{F} \quad (2)$$

171 where  $\mathbf{M}$  represents the mass matrix,  $\mathbf{C}$  is the damping matrix,  $\mathbf{K}$  denotes the stiffness matrix,  
172  $\mathbf{F}$  is the external force vector,  $\mathbf{u}$ ,  $\dot{\mathbf{u}}$  and  $\ddot{\mathbf{u}}$  correspond to the displacement, velocity, and  
173 acceleration vectors, respectively.

174 In the dynamic analysis, Rayleigh damping is applied to the entire structure and is defined as:

$$175 \quad \mathbf{C} = \alpha\mathbf{M} + \beta\mathbf{K} \quad (3)$$

176 where  $\alpha$  and  $\beta$  are the mass and stiffness coefficients, respectively, determined using the first  
177 and second natural frequencies of the structure. The Rayleigh damping ratio for the offshore jacket

178 platform is set to 5% in addition a 2% for hydrodynamic damping and 3% for structural damping  
 179 (Ajamy et al., 2018; Xu et al., 2025a).

180 To account for fluid-structure interaction during seismic events, the added mass method is  
 181 employed. The effective mass ( $m_e$ ) of the structure submerged in seawater is calculated as the sum  
 182 of the physical mass ( $m_p$ ) and the added mass ( $m_a$ ):

$$183 \quad m_e = m_p + m_a \quad (4)$$

$$184 \quad m_a = C_a \rho_w A_p \quad (5)$$

185 where  $A_p$  is the cross-sectional area of the structural member,  $\rho_w$  is the density of  
 186 seawater (taken as 1030 kg/m<sup>3</sup>), and  $C_a$  is the added mass coefficient and is set to 1.0 according  
 187 to DNV-RP-C205. Consequently, the effective density of the steel structural members submerged  
 188 in seawater is determined to be 8880 kg/m<sup>3</sup> (Xu et al., 2025c).

189 **4. Results and Discussions**

190 **4.1 Characteristics of microseismic noise power**

191 The polarization magnitudes obtained within individual weak-polarization noise sub-windows  
 192 are not entirely stable (Koper and Hawley, 2010). In addition, low-probability transient signals such  
 193 as seismic events, system calibrations, and environmental disturbances can also influence the  
 194 characteristics of microseismic noise derived from continuous recording data. Therefore, to capture  
 195 the non-random properties of the noise field at a substation, it is necessary to analyze a large number  
 196 of measurements over sufficiently long time to yield practically meaningful results. Moreover, to  
 197 mitigate the effect of random signals on the inherent microseismic noise characteristics of the  
 198 substation, we use a probability density function (PDF) approach by extracting the maximum  
 199 probability distribution curve to represent the station's intrinsic microseismic noise characteristics.

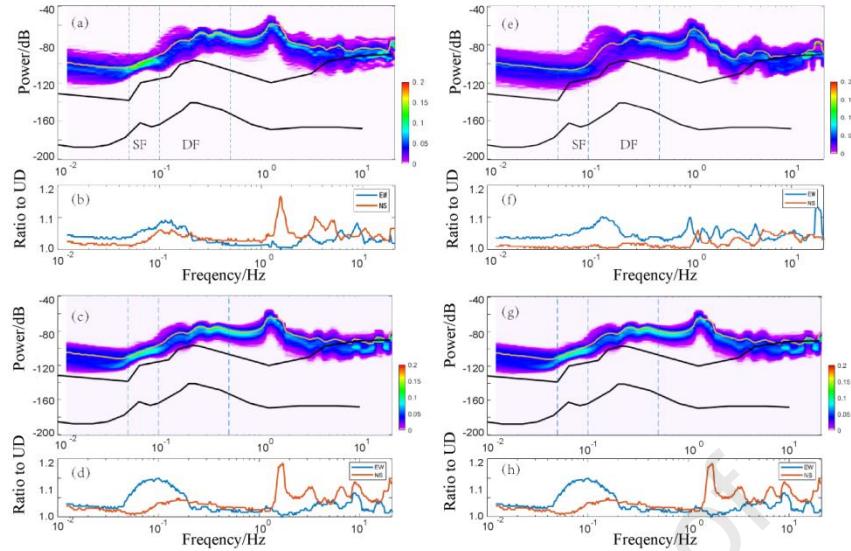
200 Fig. 4 shows the power ( $\lambda_0$ ) of microseismic noise at four offshore platform monitoring stations.  
 201 The noise power at the stations generally exceeds that observed from Peterson's (1993) High Noise  
 202 Model (NHNM) across the entire frequency range. The average noise power within the 1~20 Hz  
 203 frequency range is approximately -91 dB and is higher than the Class V ambient noise power  
 204 typically observed at land-based stations. A distinct peak is observed between 1-2 Hz, likely  
 205 attributed to the resonance effects of the platform structure. Within the microseismic frequency  
 206 band, the energy of single-frequency (SF) microseisms is clearly visible, while the double-  
 207 frequency (DF) band exhibits a segmentation phenomenon near 0.2 Hz. The spectral power is

208 predominant in the frequency ranges of approximately 0.1~0.25 Hz and 0.25~1.0 Hz, with peaks  
209 occurring near 0.2 Hz (5 s) and 0.4 Hz (2.5 s), respectively.

210 To further elucidate the noise characteristics observed at offshore platform stations, Fig. 5  
211 presents the noise power spectra from two island-based stations. These two stations are the  
212 Beihuangcheng Island Station (BHC), located in the Bohai Strait, and the Qiansandao Station  
213 (QSD), situated in the sea area of Lianyungang, Jiangsu Province. Both stations are approximately  
214 40 km from the nearest landmass, with their observation equipment deployed on bedrock. The  
215 locations of these two stations are also illustrated in Fig. 1. The noise characteristics in Fig. 5 are  
216 similar to those observed at island stations in comparable environments. The results demonstrate  
217 that the continuous seismic data recorded at offshore platform stations effectively capture the  
218 microseismic noise characteristics of island-based stations. This finding underscores the capability  
219 of offshore platforms to provide reliable and representative noise data, consistent with observations  
220 from nearby island stations.

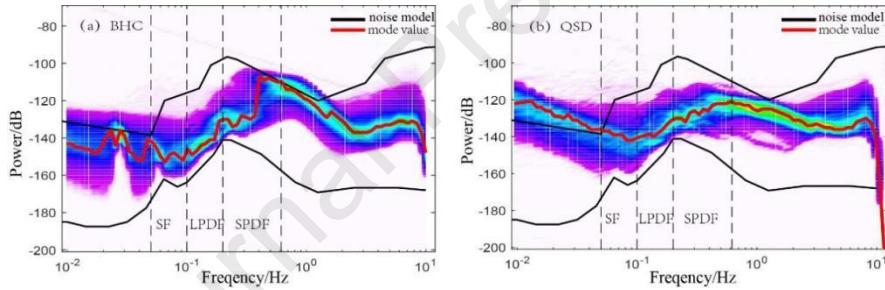
221 Moreover, according to the previous studies (Bromirski et al., 2005; Sun et al., 2013; Koper  
222 and Burlacu, 2015), DF microseisms are classified into long-period double-frequency (LPDF) and  
223 short-period double-frequency (SPDF) microseisms, with the energy of SPDF being stronger than  
224 that of LPDF. Bromirski et al. (2013) suggested that near-coastal reflections are the primary source  
225 for microseismic noise above 0.12 Hz, primarily generated by wave interactions involving gravity  
226 waves with periods of approximately 18 seconds (wavelengths of approximately 500 m) or shorter.  
227 Spectral peaks above 0.2 Hz are attributed to a relatively large source region formed by the  
228 interaction of shorter-period monsoon-forced waves near the station. The wave period in the Yellow  
229 Sea ranges between 3-3.4 seconds, which theoretically can generate a frequency spectrum above  
230 0.2 Hz. This explains why distinct segmentation phenomena are clearly observed near 0.2 Hz within  
231 the DF frequency band range (see Fig. 4).

232



233

234 Fig. 4 Polarization analysis from four offshore platform monitoring stations: the power of ambient seismic noise for  
 235 (a) HYB05, (c) HYB06, (e) HYB07, and (g) HYS02; the ratio of the maximum horizontal power to the maximum  
 236 vertical power for (b) HYB05, (d) HYB06, (f) HYB07, and (h) HYS02



237

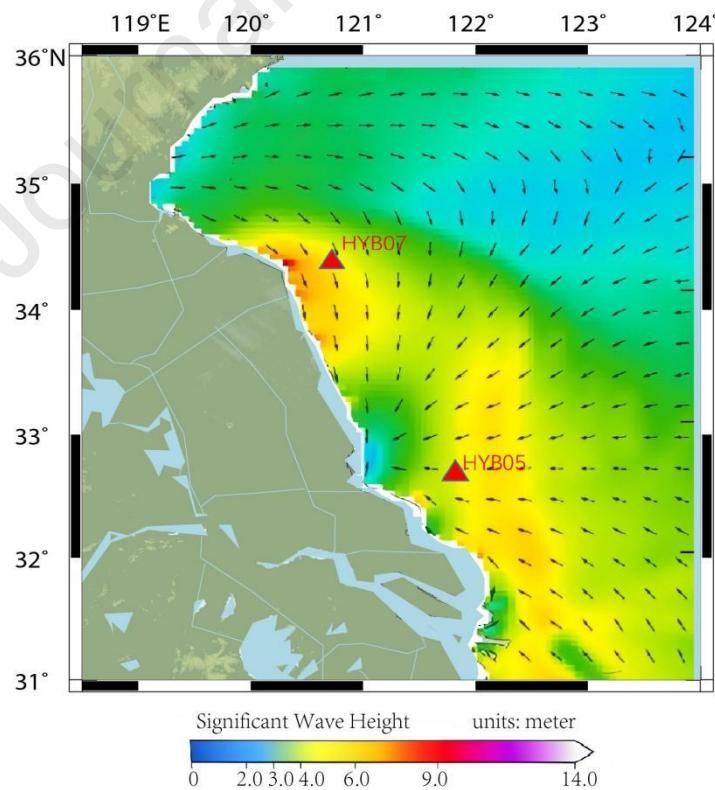
238 Fig. 5 The power of ambient seismic noise at island-based stations: (a) BHC and (b) QSD

239 By comparing Fig. 4 and 5, it is evident that the noise power at the offshore platform  
 240 monitoring stations exhibit an overall structural amplification effect compared to those of the  
 241 island-based stations. In addition, the ratio of the maximum horizontal power to the maximum  
 242 vertical power consistently exceeds 1 across all frequencies (Fig. 4b, 4d, 4f, and 4h). This indicates  
 243 that the amplification effect of microseismic noise energy is stronger in the two horizontal  
 244 components than in the vertical component. A peak amplification effect with a normal distribution  
 245 shape is observed in the low-frequency range near 0.1 Hz, while the amplification fluctuates in the  
 246 high-frequency range above 1 Hz and is strongest near 1.5 Hz.

247 Fig. 4b also shows that the amplification effect in the NS component is stronger than that in  
 248 the EW component at the HYB05 station. In contrast, the amplification effect in the EW component  
 249 is stronger than that in the NS component at the HYB07 station (Fig. 4f). The amplification effect  
 250 of ambient seismic noise is more pronounced in the EW component at approximately 0.1 Hz for

251 both HYB06 and HYS02 stations. This is because of specific environmental factors such as  
 252 hydrology, sea conditions, wind, and waves at the platform locations, which can enhance noise  
 253 energy at specific frequencies and directions (Webb, 1998; Tsai, 2011).

254 In the Yellow Sea, the sea waves are predominantly wind-generated waves, with the effective  
 255 wave height being significantly greater than that of swell. Moreover, northward waves account for  
 256 55% of the total wave frequency. In the northern Yellow Sea, northwest waves rank second in  
 257 frequency, while in the central and southern Yellow Sea, north-northeast waves are secondary, both  
 258 with frequencies ranging from 20% to 25% (Chen et al., 2016). According to data from the National  
 259 Marine Environmental Forecasting Center, the average wave direction in the Yellow Sea between  
 260 January 2024 and December 2024 is plotted in Fig. 6. Overall, the wave direction in the area of  
 261 station HYB05 is predominantly EW direction, with the maximum microseismic noise power in the  
 262 NS direction being stronger than that in the EW direction. The microseismic noise observed from  
 263 the HYB07 station exhibits similar characteristics, with waves approaching a NS direction, but the  
 264 maximum microseismic noise power in the EW direction surpasses that in the NS direction. In other  
 265 words, waves amplify microseismic noise power perpendicular to their propagation direction.  
 266

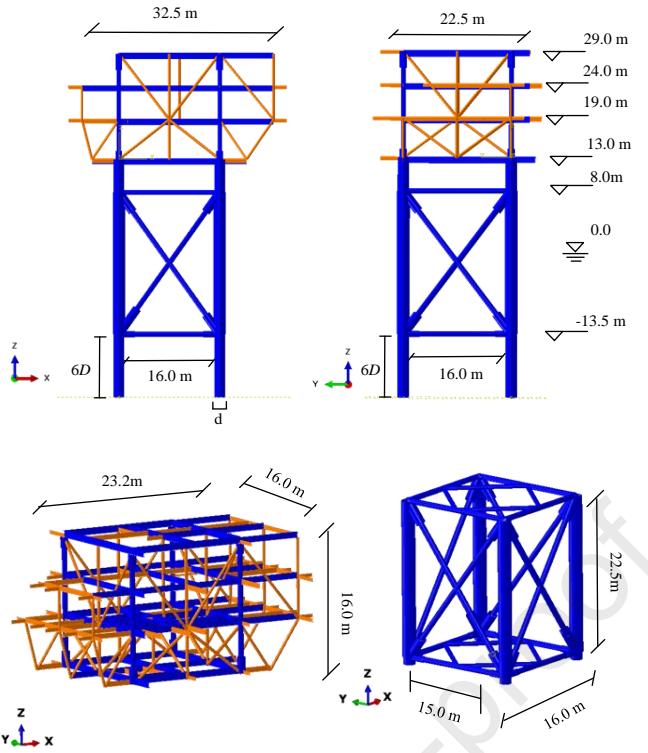


267  
 268 Fig. 6 Wave direction distribution in the Yellow Sea area between January 2024 and December 2024

269 **4.2 Correlation between microseisms and platform structure response**

270 As discussed in Section 3.1, the ambient seismic noise recorded at the offshore platform  
271 monitoring stations demonstrates a general amplification effect compared to that from the land-  
272 based stations for two primary reasons: (1) The amplitude of the Earth's ambient seismic noise  
273 significantly amplifies as it propagates through the seabed mudline to the first floor of the offshore  
274 platform, covering a distance of 26.5 m. According to the Chinese GB 50011-2010 and the  
275 American ASCE 7 standards, the ratio of the peak floor acceleration of steel structures to the base  
276 input peak acceleration (i.e., the peak floor amplification ratio) can reach up to 3. Research by Zou  
277 et al. (2023) indicates the floor amplification factor of steel structures may even exceed 5; (2) The  
278 offshore wind power platform structure generates dynamic responses under the action of waves,  
279 wind, and other environmental factors, increasing the likelihood of resonance and subsequently  
280 amplifying the surrounding seismic noise.

281 The observations demonstrate that recorded data from offshore platform monitoring stations  
282 exhibit significant amplification effects, which are closely related to the structural modal  
283 characteristics. This study conducted structural modal analysis and dynamic response analysis of a  
284 typical offshore platform to quantitatively investigate the amplification effects. This can provide a  
285 scientific basis for properly utilizing platform monitoring station data for earthquake rapid reporting,  
286 earthquake early warning, and noise seismology studies. To further quantify the amplification factor  
287 of the offshore structure on the ambient seismic noise, this study performed finite element analysis  
288 of the dynamic structural response under various frequency excitations. We chose a typical  
289 monitored platform (i.e., HYS02) for seismic response analysis, which is composed of an upper  
290 four-floor platform, a lower jacket foundation, and 4 steel pipe piles (see Fig. 7). The platform's  
291 beams are primarily made up of I-beams. In finite element analysis, the substation structure is  
292 modeled using beam elements (B31), with materials adopting an ideal elastic-plastic constitutive  
293 model. Fig. 7 also shows that the structure mainly utilizes two types of materials: Q355C for the  
294 orange components and DH36 for the blue components, with specific parameters detailed in Table  
295 2.



296  
297 Fig. 7 Finite element modeling of the monitored HYS0 substation structure  
298  
299

Table 2 Material parameters

Material type	Density (kg/m <sup>3</sup> )	Elastic parameter		Plastic parameter
		Yong's Modulus (Pa)	Poisson's Ratio	Yield Stress (Pa)
DH36	7850	$2.1 \times 10^{11}$	0.3	$3 \times 10^8$
Q355C	7850	$2.06 \times 10^{11}$	0.3	$3.55 \times 10^8$

300 To account for the effect of pile-soil interaction on the upper structure, the numerical model is  
301 fixed at a location 6 times the pile diameter below the base of the jacket (Ye and Xin, 2024), and  
302 modal analysis is carried out on the offshore structure. Table 3 gives the first 5 modes of the offshore  
303 structure.

304 Table 3 First 5 modes of the monitored HYS0 offshore substation structure

Mode	Number Frequency (Hz)	Circular Frequency (rad/s)
1	1.1991	7.534168
2	1.459	9.167167
3	1.5903	9.99215
4	1.8402	11.56232
5	2.4311	15.27505

305  
306 In the dynamic analysis, a sinusoidal acceleration excitation with an amplitude of 1 m/s is  
307 introduced at the fixed base. Fig. 8 shows the relationship between the peak floor amplification

ratio and the frequency of input sinusoidal waves at four floors of the substation structure. The results show that the peak floor amplification ratios exhibit a similar pattern with the input frequency across different floors and generally increase with the floor's height. As a results, the peak floor amplification ratios peak at the top floor. Notably, the peak floor amplification ratios peaks between 1.2 Hz and 1.6 Hz with corresponding values ranging from 6.6 to 7.7. This can be attributed to the excitation frequency being close to the first three modal frequencies of the offshore structure, thereby causing the resonance of the structure. This can support the observed peak amplitudes in the intensity of ambient seismic noise within the range of 1~2 Hz shown in Fig. 4a, 4c, 4e, and 4g.

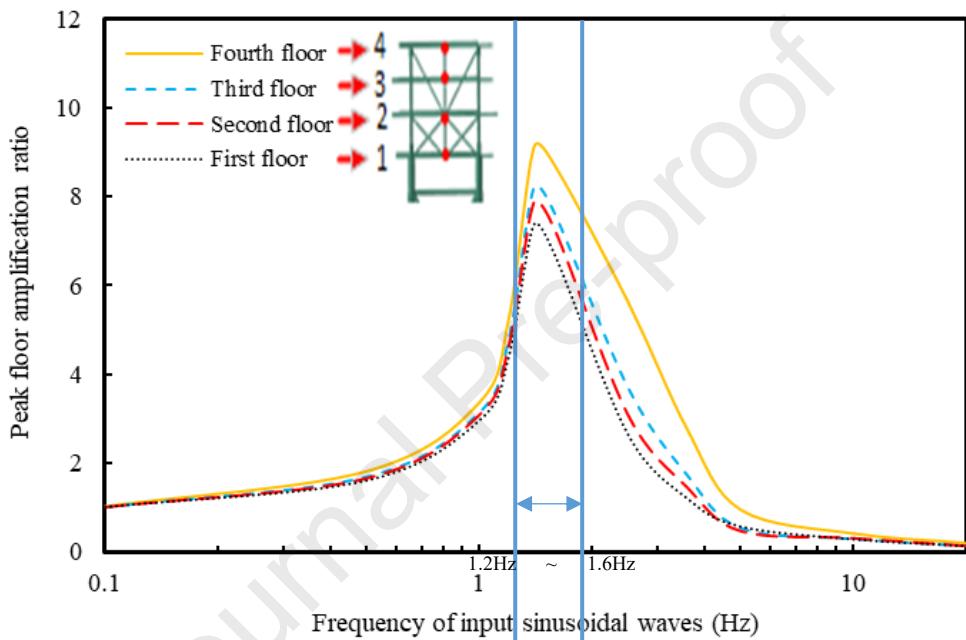


Fig. 8 Relationship between the peak floor amplification ratio and the frequency of input sinusoidal waves at four floors of the substation structure.

### 4.3 Analysis of Microseismic Polarization Characteristics

To further analyze the characteristics of microseismic noise at the offshore platform monitoring stations, Fig. 9a and 9e present the  $\theta_H$  obtained from the noise polarization analysis. It can be observed that the azimuth of the primary polarization axis is predominantly concentrated around  $200^\circ$  within the microseismic frequency band. This is consistent with the findings of Wang et al. (2022), who reported that the polarization azimuths of most stations in eastern China are mainly distributed between  $160^\circ$  and  $240^\circ$ . This suggests that the noise sources recorded at the offshore platform stations are similar to those observed at land-based stations, with the dominant polarization direction projected along a great-circle path pointing toward the Indian Ocean in the southwest (Wang et al., 2022). For noise signals above 1 Hz, the polarization directions exhibit

330 higher clustering across the entire frequency range. The  $\theta_H$  at the HYB05 station shows a primary  
 331 distribution along  $30^\circ$  to  $200^\circ$ , while the  $\theta_H$  at the HYB07 station varies between  $150^\circ$  and  $350^\circ$ .  
 332 The high-frequency signals primarily originate from near-field noise sources, which may be  
 333 associated with the inherent vibration sources of offshore wind turbine platform facilities. The  
 334 variation in signal directionality correlates with the relative installation positions of the monitoring  
 335 equipment on the platform.

336 For the low-frequency band below 0.05 Hz, the  $\beta^2$  exhibits a scattered distribution ranging  
 337 from 0.2 to 0.8 (see Fig. 9b and 9f). This indicates that the SF microseisms show insignificant  
 338 kinematic differences between vertical and tangential motions. The dominant eigenvector does not  
 339 represent pure Rayleigh waves but likely a mixture of Rayleigh waves, Love waves, and body  
 340 waves, consistent with observations from land-based stations in low-frequency ranges (Koper and  
 341 Burlacu, 2015). Moreover, a distinct and pronounced peak of approximately 0.9 near 0.1 Hz is  
 342 observed within the microseismic frequency band. This is followed by a decrease to approximately  
 343 0.2 as the frequency increases toward 0.2 Hz. Correspondingly, the  $\theta_V$  exhibits clear segmentation  
 344 characteristics at 0.1 Hz, with relatively broad  $\theta_V$  distributions observed around this frequency (see  
 345 Fig. 9c and 9g).

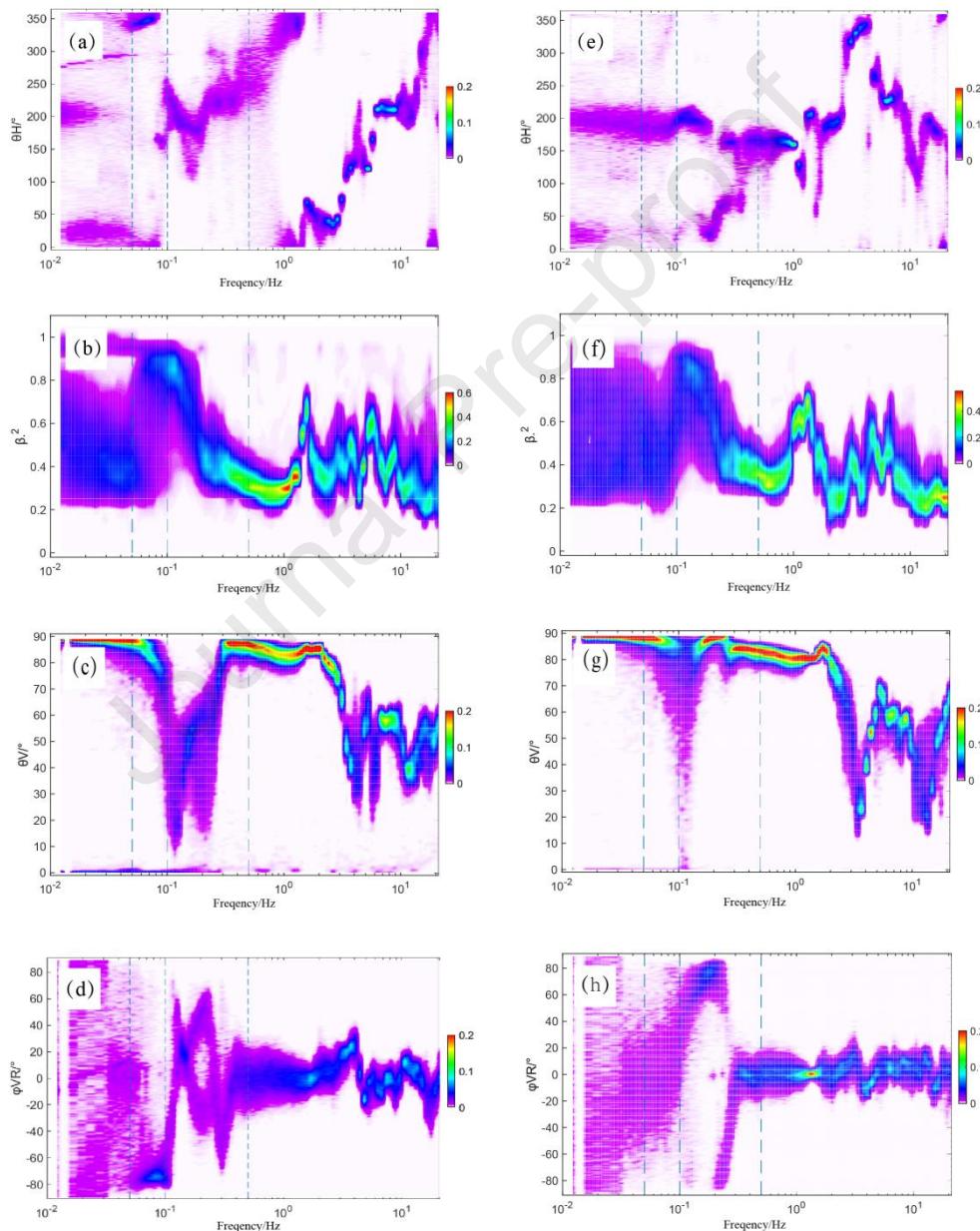
346 Fig. 9d and 9h shows that the  $\varphi_{VR}$  probability is widely distributed in the microseismic band,  
 347 further suggesting weak vertical motion. However, the  $\theta_V$  probability on both sides peak sharply  
 348 near  $90^\circ$  (see Fig. 9c and 9g). The results indicate that horizontal motion predominates over vertical  
 349 motion in both the 12-20s and 2-10s frequency bands. This implies a notable platform resonance  
 350 effect at 0.1 Hz that amplifies vertical microseismic signals.

351 For frequencies above 1 Hz, the  $\beta^2$  fluctuates with increasing frequency for both offshore  
 352 stations and maintains intermediate levels between 0.3 and 0.6. In contrast, land-based stations  
 353 typically show lower polarization degrees ( $\beta^2 < 0.2$ ) in this high-frequency band. This indicates that  
 354 stronger polarization can be observed in offshore platform recordings at certain frequencies,  
 355 especially revealing significant differences in high-frequency wavefield characteristics between  
 356 marine and terrestrial stations. These differences are likely due to short-period waves generated by  
 357 wind-wave interactions in the sea, where wave breaking and turbulent processes can generate high-  
 358 frequency noise.

359 The corresponding vertical-radial phase difference ( $\varphi_{VR}$ ) follows a near-zero degree  
 360 distribution, indicating synchronous propagation of vertical and radial components of seismic wave

361 with zero phase difference between two components. This phenomenon typically occurs during  
 362 compressional (P-) wave propagation, where the vibration direction aligns with the propagation  
 363 direction, representing simple and symmetric wave modes. Such observations may also suggest  
 364 either homogeneous medium properties or relatively simple wave propagation paths. Moreover, for  
 365 frequencies above 1 Hz, approximately zero  $\varphi_{VR}$  implies simple wave propagation patterns, with  
 366 geometric configurations or polarization states exhibiting high degrees of symmetry.

367



368

369 Fig. 9 Polarization analysis results for two stations:  $\theta_H$  for (a) HYB05 and (e) HYB07,  $\beta^2$  for (b) HYB05 and (f)  
 370 HYB07,  $\theta_V$  for (c) HYB05 and (g) HYB07, and  $\varphi_{VR}$  for (d) HYB05 and (h) HYB07  
 371 Moreover, distinct segmentation characteristics are observed near 0.2 Hz for  $\theta_V$  and  $\varphi_{VR}$  of  
 372 both stations' DF microseisms, indicating non-identical vibration properties in vertical and radial

373 components. This suggests different energy generation mechanisms and distinct field sources  
 374 between the two stations. The long-period double-frequency (LPDF) demonstrates stronger vertical  
 375 vibration vectors with relatively dominant Rayleigh wave components, while the short-period  
 376 double-frequency (SPDF) shows comparatively more Love wave components. These  
 377 characteristics resemble observations from land-based permanent stations, confirming the  
 378 reliability of at least two persistent field sources for DF microseisms in eastern China. The LPDF  
 379 microseisms are more likely generated through wave-seafloor interactions in open coastal waters  
 380 (Xiao and Huang, 2015; Xiao et al., 2018). However, the higher  $\beta^2$  observed at offshore platforms  
 381 compared to land stations may reflect non-negligible direct impacts from marine hurricanes/storm  
 382 activities and corresponding platform responses.

## 383 5. Conclusions

384 This study adopts frequency-domain polarization analysis to investigate the ambient seismic  
 385 noise characteristics (energy intensity, wavefield composition, and polarization properties) using  
 386 continuous waveform data from seismometers deployed on offshore wind platforms.  
 387 Complementary finite element modeling is employed to quantitatively assess platform structural  
 388 responses to ambient noise excitation. The main conclusions are as follows:

- 389 (1) The average noise power at offshore platform stations ranges around -91 dB within the 1-  
 390 20 Hz frequency band. Distinct SF energy is observed in the microseismic band, while  
 391 the DF band exhibits clear segmentation near 0.2 Hz. This feature is consistent with that  
 392 observed from island and coastal stations. Horizontal components show higher energy  
 393 than vertical components, likely due to lateral wave/current forces amplifying platform  
 394 vibrations directionally.
- 395 (2) Compared to land stations, offshore platforms demonstrate the noise amplification in the  
 396 whole frequency range. Notably, the peak floor amplification ratios peaks between 1.2 Hz  
 397 and 1.6 Hz with corresponding values ranging from 6.6 to 7.7. This can be attributed to  
 398 the excitation frequency being close to the first three modal frequencies of the offshore  
 399 structure, thereby causing the resonance of the structure. This can support the observed  
 400 peak amplitudes in the intensity of ambient seismic noise within the range of 1~2 Hz.
- 401 (3) The continuous recording data from offshore engineering platform monitoring stations  
 402 can be utilized for studying the properties of microseismic wavefields. Within the  
 403 microseismic frequency band (20s-0.5Hz), the azimuth of the polarization principal axis  
 404 of noise is predominantly concentrated around 200°. The low-frequency components  
 405 (<0.05 Hz) exhibit a scattered distribution of  $\beta^2$  ranging between 0.2 and 1. This indicates  
 406 negligible kinematic differences between vertical and tangential motions of SF

407 microseisms, where the principal eigenvector does not represent pure Rayleigh waves but  
 408 rather a hybrid composition of Rayleigh, Love, and body waves. For the high-frequency  
 409 band (>1 Hz), the polarization degree exhibits perturbed variations within the 0.2–0.6  
 410 range as frequency increases, with  $\beta_2$  values systematically higher than those observed at  
 411 onshore stations..

412 Due to the high noise levels, offshore platforms might not be ideal locations for seismic  
 413 observations. This study employs finite element modeling to quantitatively analyze the  
 414 amplification effects of platform structures on ambient seismic noise, providing important  
 415 theoretical references for subsequently improving the signal-to-noise ratio of observational  
 416 data. Meanwhile, these monitoring data can be further utilized for safety assessments of the  
 417 offshore structure in the future.

#### 418 **Acknowledgements**

419 This research was jointly supported by the Spark Program Project of the China Earthquake  
 420 Administration (XH23015A), the Industry Foresight and Key Core Technology Project of the  
 421 Jiangsu Province Key Research and Development Program (BE2023082), the Director Fund  
 422 Project of the Administration of Jiangsu Earthquake (202401), and the National Natural Science  
 423 Foundation of China (52378345). We thank two anonymous reviewers, the associate editor and  
 424 deputy editor-in-chief for their value comments during the review process.

#### 425 **6. References**

- 426 Ajamy, A., Asgarian, B., Ventura, C., Zolfaghari, M., 2018. Seismic fragility analysis of jacket type offshore platforms  
 427 considering soil-pile-structure interaction. *Engineering Structures*, 174: 198-211.
- 428 American Society of Civil Engineers, ASCE 7-16: Minimum Design Loads for Buildings and Other Structures, 2016.  
 429 Reston, Virginia
- 430 Baba, S., Araki, E., Yokobiki, T., Kawamata, K., Uchiyama, K., Yoshizuka, T., 2024. Seismic observation using  
 431 distributed acoustic sensing around the Tsugaru Strait at the Japan and Kuril Trenches, northeastern Japan.  
 432 *Earth Planets and Space*, 76(1).
- 433 Bromirski, P.D., Duennebier, F.K., Stephen, R.A., 2005. Mid-ocean microseisms. *Geochemistry, Geophysics,  
 434 Geosystems*, 6(4).
- 435 Bromirski, P.D., Stephen, R.A., Gerstoft, P., 2013. Are deep-ocean-generated surface-wave microseisms observed on  
 436 land? *Journal of Geophysical Research: Solid Earth*, 118(7): 3610-3629.
- 437 Chen, G., Qu, F., Li, P., Liu, Xiao., 2016. Numerical study of wave height seasonality in the Yellow Sea, *Marine  
 438 Sciences*, 40(11): 155-168.
- 439 Chen, W., Luo, Z., Zheng, Y., Xu, L., & Huang, L. (2025). Seismic response analysis of liquefiable seabed sites: Effects  
 440 of seawater-seabed interaction and bidirectional seismic motion. *Ocean Engineering*, 324, 120717.
- 441 Chen, Y., Gao, L., Zhao, F., 2007. A method to enhance the signal/noise ratio of three-component seismic data based

- 442 on the polarization analysis in frequency domain. *Progress in Geophysics*, 2007, 22(1): 255-261.
- 443 Flinn, E.A., 1965. Confidence regions and error determinations for seismic event location. *Reviews of Geophysics*,  
444 3(1): 157-185.
- 445 GB 50011-2010. *Code for Seismic Design of Buildings*. Beijing: China Architecture and Building Press; 2010. (In  
446 Chinese)
- 447 Gong, J., Sun, Y., Zheng, J., 2023. Construction and Efficiency Analysis of Jiangsu Marine Seismic Stations. *Journal*  
448 *of Disaster Prevention and Mitigation Engineering*, 43(03):614-625.
- 449 Hasselmann, K., 1963. A statistical analysis of the generation of microseisms. *Reviews of Geophysics*, 1(2): 177-210.
- 450 Ito, A., Yamamoto, Y., Hino, R., Suetsugu, D., Sugioka, H., Nakano, M., Obama, K., Nakahigashi, K..Shinohara,  
451 M., 2017. Tomographic image of crust and upper mantle off the Boso Peninsula using data from an ocean-  
452 bottom seismograph array. *Earth Planets and Space*, 69.
- 453 Jepsen, D..Kennett, B., 1990. Three-component analysis of regional seismograms. *Bulletin of the Seismological*  
454 *Society of America*, 80(6B): 2032-2052.
- 455 Jurkevics, A., 1988. Polarization analysis of three-component array data. *Bulletin of the Seismological Society of*  
456 *America*, 78(5): 1725–1743.
- 457 Koper, K.D..Burlacu, R., 2015. The fine structure of double-frequency microseisms recorded by seismometers in North  
458 America. *Journal of Geophysical Research: Solid Earth*, 120(3): 1677-1691.
- 459 Koper, K.D..Hawley, V.L., 2010. Frequency dependent polarization analysis of ambient seismic noise recorded at a  
460 broadband seismometer in the central United States. *Earthquake Science*, 23: 439-447.
- 461 Krylov, A.A., Egorov, I.V., Kovachev, S.A., Ilinskiy, D.A., Ganzha, O.Y., Timashkevich, G.K., Roginskiy, K.A.,  
462 Kulikov, M.E., Novikov, M.A., Ivanov, V.N., Radiuk, E.A., Rukavishnikova, D.D., Neeshpapa, A.V.,  
463 Velichko, G.O., Lobkovsky, L.I., Medvedev, I.P..Semiletov, I.P., 2021. Ocean-Bottom Seismographs Based  
464 on Broadband MET Sensors: Architecture and Deployment Case Study in the Arctic. *Sensors*, 21(12).
- 465 Lin, C.-r., Liao, Y.-C., Wang, C.-c., Kuo, B.-Y., Chen, H.-H., Jang, J.-p., Chen, P.-c., Chang, H.-k., Lin, F.-S..Chang,  
466 K.-H., 2024. Development and evaluations of the broadband ocean bottom seismometer (Yardbird-BB OBS)  
467 in Taiwan. *Terrestrial, Atmospheric and Oceanic Sciences*, 35(1): 4.
- 468 Lin, J., Yang, W., Chen, Meng., Wu, Renhao., Ge Hongkui., 2012. The application of polarization analysis in seismic  
469 signal detection. *Earthquake Research In China*, 28(2): 133-143.
- 470 Miche, A., 1944. Mouvements ondulatoires de la mer en profondeur constante ou décroissante. *Ann. Ponts Chaussees*,  
471 25: 369-406.
- 472 Park, J., Vernon III, F.L..Lindberg, C.R., 1987. Frequency dependent polarization analysis of high-frequency  
473 seismograms. *Journal of Geophysical Research: Solid Earth*, 92(B12): 12664-12674.
- 474 Peterson, J.R., 1993. Observations and modeling of seismic background noise. 2331-1258, US Geological Survey.
- 475 Samson, J., 1983. Pure states, polarized waves, and principal components in the spectra of multiple, geophysical time-  
476 series. *Geophysical Journal International*, 72(3): 647-664.
- 477 Shinohara, M., Yamada, T., Akuhara, T., Mochizuki, K..Sakai, S., 2022. Performance of Seismic Observation by  
478 Distributed Acoustic Sensing Technology Using a Seafloor Cable Off Sanriku, Japan. *Frontiers in Marine*  
479 *Science*, 9.
- 480 Stephen, R., 2003. Workshop on Seismo-Acoustics Applications in Marine Geology and Geophysics. National Science  
481 Foundation.
- 482 Sun, T., Xue, M., Le, K.P., Zhang, Y..Xu, H., 2013. Signatures of ocean storms on seismic records in South China  
483 Sea and East China Sea. *Marine Geophysical Research*, 34(3-4): 431-448.
- 484 Sun, Y., He, Y., Zheng, J., Wang, J.. 2022. A Preliminary Investigation of an M3.8 Earthquake in the South Yellow  
485 Sea Recorded by an Experimental Offshore Wind Power Platform based Seismic Station . *Journal of Disaster*  
486 *Prevention and Mitigation Engineering*,(6):1351-1353
- 487 Takagi, R., Nishida, K., Maeda, T..Obara, K., 2018. Ambient seismic noise wavefield in Japan characterized by

- 488 polarization analysis of Hi-net records. *Geophysical Journal International*, 215(3): 1682-1699.
- 489 Tanimoto, T..Prindle, K., 2007. Surface wave analysis with beamforming. *Earth, planets and space*, 59: 453-458.
- 490 Tsai, V.C., 2011. Understanding the amplitudes of noise correlation measurements. *Journal of Geophysical Research: Solid Earth*, 116(B9).
- 491 Viana P. B., Kiwamu N., Bogdan E., Ryota T.; 2025. Ocean Microseisms Recorded by the Cuban Seismic Network: Time Variation, Spectral Features, and Source Directionality Characteristics. *Seismological Research Letters*; 494 96 (2A): 758-771
- 495 Wagner, G.S..Owens, T.J., 1996. Signal detection using multi-channel seismic data. *Bulletin of the Seismological 496 Society of America*, 86(1A): 221-231.
- 497 WANG Jun, ZHENG DingChang, ZHANG JinChuan, ZHAN XiaoYan, Qian Ting. 2020. Seismic velocity changes in 498 the epicentral region of the 2013 Lushan earthquake measured from ambient seismic noise. *Chinese Journal 499 of Geophysics* (in Chinese), 63(2): 517-531
- 500 WANG Jun, ZHENG XiuFen, LIU HongGui, YANG DaKe, ZHENG DingChang, ZHENG JiangRong, ZHAN XiaoYan, 501 SUN YeJun, DU Hang. 2022. Characteristics and location of non-random persistent microseismic ambient 502 noise sources in eastern China. *Chinese Journal of Geophysics* (in Chinese), 65(5): 1660-1674
- 503 Webb, S.C., 1998. Broadband seismology and noise under the ocean. *Reviews of Geophysics*, 36(1): 105-142.
- 504 Xiao, H.., Xue, M.., Pan, M..Gao, J., 2018. Characteristics of microseisms in South China. *Bulletin of the Seismological 505 Society of America*, 108(5A): 2713-2723.
- 506 Xiao, H..Huang, W., 2015. Three-dimensional numerical modeling of solitary wave breaking and force on a cylinder 507 pile in a coastal surf zone. *Journal of Engineering Mechanics*, 141(8): A4014001.
- 508 Xu, L. Y., Huang, J. T., Liu, L., Song, C. X., Zhou, Z. L., Li, Z. Y., Cai, F, & Chen, G. X. (2025a). Offshore wind turbine 509 monopiles under real wind and wave conditions: Considering cyclic softening of clays. *Ocean 510 Engineering*, 322, 120449. <https://doi.org/10.1016/j.oceaneng.2025.120449>
- 511 Xu, L. Y., Xi, J. P., Jiang, J. W., Cai, F., Sun, Y. J., & Chen, G. X. (2025b). Seismic fragility analysis of shield tunnels 512 in liquefiable layered deposits. *Soil Dynamics and Earthquake Engineering*, 191, 109246. 513 <https://doi.org/10.1016/j.soildyn.2025.109246>
- 514 Xu, L. Y., Yuan, Z. H., Peng, X. B., Chen, W. Y., Cai, F., Zhou, Y. G., & Chen, G. X. (2025c). Seismic response analysis 515 of offshore wind turbines supported by monopiles on gently sloping sandy seabed. *Ocean Engineering*, 323, 516 120599.
- 517 YE, S.-h..XIN, L.-l., 2024. Settlement and bearing capacity of single pile based on shear characteristics of pile-soil 518 interface. *Rock and Soil Mechanics*, 45(5): 7.
- 519 ZHU Feng, YANG Chi, HE Yicheng, et al. Preliminary discussion on the earthquake early warning network of Jiangsu 520 Province[J]. *Seismological and Geomagnetic Observation and Research*, 2023, 44(5): 52-58.
- 521 Zou, X.., Yang, W.., Liu, P..Wang, M., 2023. Floor acceleration amplification and response spectra of reinforced 522 concrete frame structure based on shaking table tests and numerical study. *Archives of Civil and Mechanical 523 Engineering*, 23(3): 156.
- 524
- 525

# Characteristics of Ambient Seismic Noise Recorded at Offshore Wind Turbine Platform Monitoring Stations

Jun Wang , Jiang-rong Zheng, Ye-jun Sun, Ling-yu Xu, Yi-cheng He , Jie Gong , Chen Wang, Xiao-yan Zhan, Yu-jie Wan , Wen-jie Ren

**1、Jun Wang**, Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.

Email: wangjun1099@qq.com, Phone: 13951702501.

**2、Jiang-rong Zheng**, Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.

Email: jsdzjzjr@126.com.

**3、Ye-jun Sun**, Professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.

Email: syj\_4116@126.com.

**4、Ling-yu Xu**, Professor Institute of Geotechnical Engineering, Nanjing Tech University, Nanjing 210009, China.

Email: lyxu@njtech.edu.cn

**5、Yi-cheng He**, Associate professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.

Email: hyckevan@mail.ustc.edu.cn

**6、Jie Gong**, Associate professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China. Email: 498602259@qq.com

**7、Chen Wang**, Engineer, Earthquake Administration of Jiangsu Province, Nanjing 210014, China.

Email: 372766770@qq.com

**8、Xiao-yan Zhan**, Associate professor, Earthquake Administration of Jiangsu Province, Nanjing 210014, China. Email: 43493270@qq.com

**9、Yu-jie Wan**, Engineer, Earthquake Administration of Qinghai Province, Xining 810000, China.

Email: 376633212@qq.com

**10、Wen-jie Ren**, PhD candidate, Institute of Geotechnical Engineering, Nanjing Tech University, Nanjing 210009, China.

Email: renwenjie@njtech.edu.cn

**Corresponding author:** Jun Wang, wangjun1099@qq.com

**Table 2** Detailed parameters of two equipment used for marine seismic observation

Serial Number	Equipment	Seismometer Sensitivity	Data Logger Conversion Factor	Sampling Rate	Frequency Bandwidth
1	GL-PCS60	1000 V/m/s	74.5nV/count	100Hz	60s-50Hz
2	GL-PS2	1000 V/m/s	74.5nV/count	100Hz	2s-50Hz

Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.